



Massive tree mortality from flood pulse disturbances in Amazonian floodplain forests: The collateral effects of hydropower production

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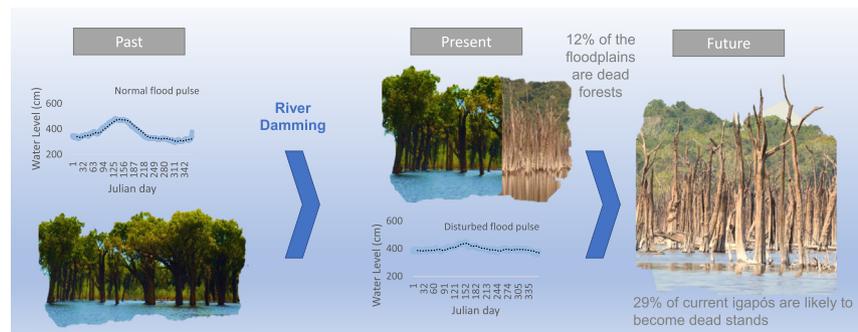
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HIGHLIGHTS

- Considerable floodplain forest loss occurred downstream of the Balbina dam from flood pulse changes;
- Several tree species are still suffering mortality, which may increase dead stand areas;
- Remote sensing methods used here may be applied to other floodplains with massive tree mortality by damming;
- Downstream igapó forests are highly sensitive to the loss of low water periods resulting from flow alteration by dam operations;

GRAPHICAL ABSTRACT



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ABSTRACT

Large dams built for hydroelectric power generation alter the hydrology of rivers, attenuating the flood pulse downstream of the dam and impacting riparian and floodplain ecosystems. The present work mapped black-water floodplain forests (*igapó*) downstream of the Balbina Reservoir, which was created between 1983 and 1987 by damming the Uatumã River in the Central Amazon basin. We apply remote sensing methods to detect tree mortality resulting from hydrological changes, based on analysis of 56 ALOS/PALSAR synthetic aperture radar images acquired at different flood levels between 2006 and 2011. Our application of object-based image analysis (OBIA) methods and the random forests supervised classification algorithm yielded an overall accuracy of 87.2%. A total of 9800 km² of *igapó* forests were mapped along the entire river downstream of the dam, but forest mortality was only observed below the first 49 km downstream, after the Morena rapids, along an 80-km river stretch. In total, 12% of the floodplain forest died within this stretch. We also detected that 29% of the remaining living *igapó* forest may be presently undergoing mortality. Furthermore, this large loss does not include the entirety of lost *igapó* forests downstream of the dam; areas which are now above current maximum flooding heights are no longer floodable and do not show on our mapping but will likely transition over time to upland forest species composition and dynamics, also characteristic of *igapó* loss. Our results show that floodplain forests

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Balbina
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Flood pulse

are extremely sensitive to long-term downstream hydrological changes and disturbances resulting from the disruption of the natural flood pulse. Brazilian hydropower regulations should require that Amazon dam operations ensure the simulation of the natural flood-pulse, despite losses in energy production, to preserve the integrity of floodplain forest ecosystems and to mitigate impacts for the riverine populations.

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1. Introduction

Hydroelectric dams are a growing threat to the extensive Amazon wetlands, having several negative environmental effects and disturbing the hydrological balance that controls the structure and function of these environments (Castello and Macedo, 2016; Foote et al., 1996; Kahn et al., 2014; Latrubesse et al., 2017; Timpe and Kaplan, 2017). Although dams are important for electricity production, they also increase greenhouse gas emissions, disrupt indigenous and riverine communities, and lead to forest degradation and losses of biodiversity as a consequence of the large areas artificially inundated by the reservoirs behind storage dams (Benchimol and Peres, 2015; Cochrane et al., 2017; Fearnside, 1995; Junk and de Mello, 1990; Kemenes et al., 2007). Most ecological studies of hydropower dam impacts assess only the inundated and surrounding areas of reservoirs, while downstream environmental alterations caused by the hydrological changes created by dams remain poorly understood (Assahira et al., 2017; Manyari and de Carvalho, 2007; Timpe and Kaplan, 2017; Williams and Wolman, 1984).

Dams typically alter the natural “flood pulse” (Junk et al., 1989), negatively affecting the ecological dynamics of downstream floodplain ecosystems (Benchimol and Peres, 2015; Fearnside, 1995; Junk and Mello, 1990; Kemenes et al., 2007). Under natural conditions, the flood pulse of large rivers is predictable, characterized by a monomodal flood pulse with an annual cycle of rising, high, receding and low water phases, which can span more than 10 m of amplitude in the case of Central Amazonia (Junk et al., 2011; Schöngart and Junk, 2007). This natural water level oscillation floods over 750,000 km² annually in the Amazon Basin (Wittmann and Junk, 2017), varying in amplitude and duration depending on the amount and seasonality of rainfall within catchments, and on river discharge and topography (Junk et al., 2011). The flood pulse is the main driver of ecological and biogeochemical processes in floodplain forests (Junk, 1989; Lewis et al., 2000), where trees display anatomical, morphological, physiological, and biochemical adaptations to survive in this seasonal environment (De Simone et al., 2003; Ferreira and Parolin, 2011; Junk, 1989; Parolin et al., 2004). For this reason, a large number of tree species are endemic to floodplain environments (Wittmann et al., 2006, 2013, 2017).

Within the Amazonian floodplains, *igapós* are forests influenced by black or clear water (oligotrophic) rivers with catchments in the Precambrian Guiana shield and Central Brazilian archaic shield, receiving much lower sediment loads than the eutrophic *várzea* forests, which are inundated by sediment-rich white-water rivers with catchments in the Andean region (Junk et al., 2011; Prance, 1979; Sioli, 1984). Due to their low sediment load, black water and clear water rivers have been the main targets of hydroelectric projects across the Brazilian Amazon. Furthermore, several *igapó* areas along clear water rivers are close to the so-called “Arc of Deforestation”, the expanding southwestern agricultural frontier in the Amazon, further compounding the threats to these systems (Lees et al., 2016).

Floodplain forest trees are recognized as good bioindicators of hydrological disturbances, due to their longevity (Schöngart et al., 2005). Many species have a seasonal growth pattern closely linked to the hydrological cycle (Schöngart et al., 2002), and changes in the regularity and duration of the natural flood cycle can lead to growth reduction or interruption and even to mortality. Assahira et al. (2017) found that a substantial increase in annual minimum water levels observed for several consecutive years in the Uatumã River, caused by water released

from the Balbina Reservoir for power generation, led to the mortality of large populations of the highly flood-adapted tree species *Macrobium acaciifolium* (Benth.) Benth. (Fabaceae) (Schlüter and Furch, 1992).

The number of hydroelectric power plants in the Amazon has grown steadily in the last 30 years, and this growth is expected to continue because these ventures are considered to be important for economic development, with efforts to understand and manage their socioecological implications often rushed and inadequate (Tilt, 2012; Tilt et al., 2009). Recent studies show conflicting data on the number of hydroelectric plants in operation, under construction and planned, even for papers published within the same year (Castello et al., 2013; Latrubesse et al., 2017; Lees et al., 2016; Timpe and Kaplan, 2017). Latrubesse et al. (2017) identified 140 dams in operation or under construction and 288 planned dams in the entire Amazon region. However, Anderson et al. (2018) reported 142 dams in operation or under construction and 160 planned dams for the Andean Amazon alone, implying a considerable underestimation from the remaining studies (Anderson et al., 2018; Castello and Macedo, 2016; Finer and Jenkins, 2012; Forsberg et al., 2017).

Disruptions in connectivity are also known to compromise the physical habitat and the ecological integrity of river systems around the world, threatening endemic species and allowing the colonization of invasive ones (Ligon et al., 1995; Mumba and Thompson, 2005; Nilsson, 2000; Ward and Stanford, 1995). In the case of large-scale Amazonian hydroelectric dams, existing studies have a very limited spatial distribution, and thus the true impacts and the extent of environmental disturbances caused by them remain only partially understood (Fearnside, 2014; Junk and Mello, 1990; Wittmann and Junk, 2017). There is a pressing need to quantify and understand the processes that occur downstream of Amazonian dams, where the natural flood pulse is completely altered by the opening and closing of floodgates, thereby altering the natural seasonality of the rivers (Assahira et al., 2017; Junk and Mello, 1990).

The construction of the Balbina Dam (Amazonas, state, Brazil) between 1983 and 1987 led to an abrupt change in the Uatumã River hydrology, which has been maintained for the past 30 years, providing an ideal system to assess ecohydrological responses to damming in *igapó* environments. Therefore, in the present study we address the following questions: (1) What is the spatial distribution and extent of tree mortality in *igapó* floodplain forests in the Uatumã River, downstream of the Balbina Dam? (2) Which recommendations can be made for future evaluations of environmental impacts of future hydroelectric dams in the Amazon region with focus on downstream areas? (3) Is *igapó* vegetation loss still occurring, and if so, in which areas?

2. Methods

2.1. Study area

We studied a 300-km stretch of the Uatumã River downstream of the Balbina hydroelectric dam, up to its confluence with the Amazon River, including major tributaries (Abacate and Jatapú rivers - see Fig. S1, showing the non-disturbed Abacate River *igapó*). The Uatumã River has a catchment area of approximately 69,500 km², corresponding to nearly 1.2% of the entire Amazon Basin (Melack and Hess, 2010).

The Balbina hydroelectric plant, built between 1983 and 1987, is known worldwide as one of the largest ecological disasters in Brazilian history (Fearnside, 1989). Its damming flooded about 4437 km² (Benchimol and Peres, 2015), but due to the low slope of the relief, the 250 MW of nominal energy produced is extremely low in relation to the flooded area. Compared to the Tucuruí hydroelectric dam on the Tocantins River, whose reservoir is nearly 100 km² smaller, the Balbina hydroelectric power plant produces almost 33 times less energy (International Rivers and ECOA, 2017).

Our study area is located approximately 150 km east of the capital city of Manaus (Amazonas state), within the Uatumã Sustainable Development Reserve (RDS Uatumã) (Fig. 1). Four distinct forest types are described for this region: upland forests (*terra firme*), oligotrophic floodplain forests (*igapó*), and two white-sand open vegetation (*campinas* and *campinaranas*) (Adeney et al., 2016; Andreae et al., 2015; IDESAM, 2009; Targhetta et al., 2015). According to data acquired between 2013 and 2015 from the Amazon Tall Tower Observatory (ATTO) project, the area has an average annual temperature of 27 °C and an average annual total rainfall of 1920 mm, with August and September being the driest months.

According to the Shuttle Radar Topography Mission (SRTM) global digital elevation model, the first 22 km of the Uatumã River downstream of the dam have an average width of 200 m (Figs. 1 and 2). For the next 13 km downstream, the river and its associated floodplain widen progressively, coinciding with a change from a rectilinear to a meandering pattern. The topographic profile from the dam to the Uatumã river mouth shows a steep descending gradient, with the

steepest portions occurring up to 35 km downstream of the dam (Fig. 2), thus matching the geomorphological pattern described above.

In contrast to *várzea* habitats and most Andean rivers, the Uatumã River originates from the Guiana Shield, like most black water rivers in the middle portion of the Amazon basin (northern rivers) (Abad et al., 2013; IDESAM, 2009; Sioli, 1984). Black water rivers are stable and have slow geomorphologic and floristic dynamics, which can be associated to the presence of monodominant populations of ancient trees (Junk et al., 2015).

Since the damming of the Uatumã River by the Balbina hydroelectric plant, the annual predictable flood pulse has been modified (Assahira et al., 2017; Timpe and Kaplan, 2017). However, the only available water level monitoring station for the Uatumã River downstream of the dam is the Cachoeira Morena gauging station, operated by the Brazilian Geological Service/Brazilian National Water Agency (CPRM/ANA) (Cachoeira Morena Station - number 16100000). However, because this station sits on a terrace, near rapids, its data do not reflect the range of absolute river stage variation (see Fig. 2), and its overall measured stage amplitude (ca. 2 m) is much smaller than the observed amplitudes downriver (up to 9 m, field measurements). The lack of proper monitoring stations downriver makes it difficult to characterize the absolute magnitudes of changes in the water level regime post-damming.

Furthermore, as the hydrological regime is presently irregular, care must be taken to separate abnormal events caused by the dam from the effects of extreme climatic events. As a consequence of the strong El Niño event in 1997/98 reducing rainfall in the Central Amazon region

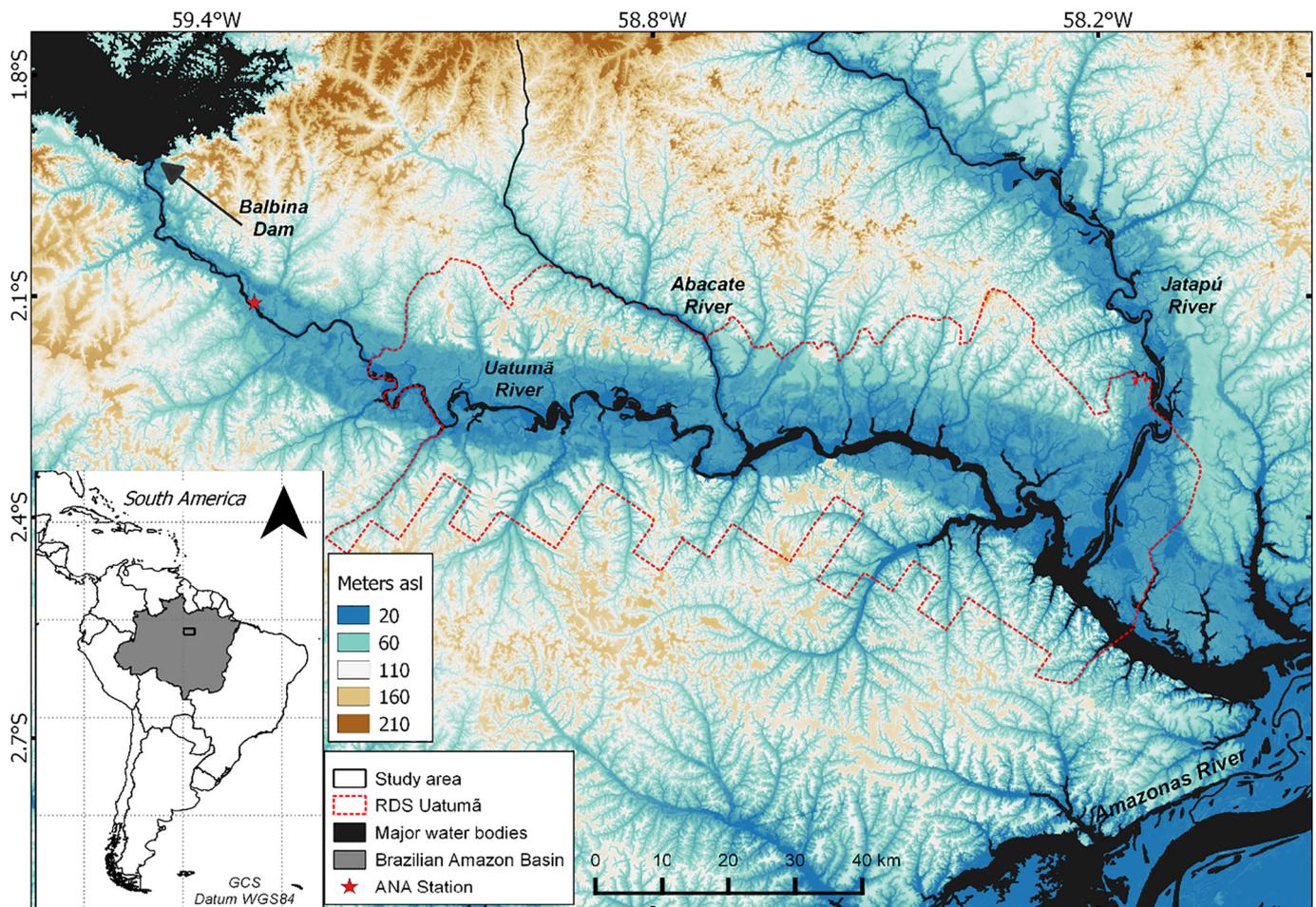


Fig. 1. Location of the study area in the Brazilian Amazon Basin, and detail of the study site, from the Balbina dam to the mouth of the Uatumã River. Background: Shuttle Radar Topography Mission digital elevation model showing the topographic variation with red outlines showing the areas mapped by remote sensing. (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

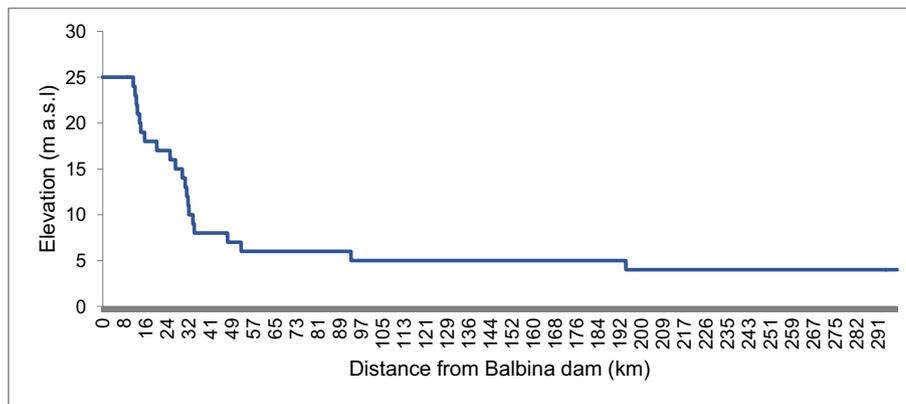


Fig. 2. Longitudinal elevation profile along the main Uatumã river channel (in February 2000) showing the steeper slope along the first 35 to 50 km downstream of the Balbina dam (SRTM data obtained from earthexplorer.usgs.gov).

(Foley et al., 2002; Marengo and Espinoza, 2016), all *igapós* were dry, followed by a 13-year period between 1999 and 2012, where all lower elevation *igapós* and some intermediate elevation *igapós* were completely flooded. In May 2000, water levels flooded all *igapós* for at least six days, reaching upland forests and riverine houses. During the severe El Niño event in 2015/16, water levels were so low that even the lowest areas including dead stands remained dry. This shows how unpredictable and irregular the flood pulse has become after starting hydropower operations, and the full consequences of such events remain unknown (Fig. 3).

2.2. Remote sensing data

We used L-band synthetic aperture radar (SAR) satellite images, as they have the ability to penetrate forest canopies and detect the flooding beneath, producing an increase in the returned signal (radar backscattering) known as “double-bounce” (Silva et al., 2015). Furthermore, L-band radar is insensitive to cloud cover and other atmospheric interferences like gases, smoke and illumination (Woodhouse, 2017). For our study, the detection of areas with high tree mortality is based on the even stronger SAR return signal caused by double-bounce between the free water surface and the standing dead trees, without the attenuation caused by volumetric scattering from canopies that is observed for live forest stands (Hess et al., 2002; Silva et al., 2008). In general, dead tree stands under flooded conditions have higher SAR returns than live flooded stands, which are still brighter than upland, non-flooded forests.

We acquired 56 SAR images from the Phased Array Type L-band Synthetic Aperture Radar (PALSAR) sensor, onboard the Advanced Land Observing Satellite-1 (ALOS-1), operated by the Japanese Aerospace Exploration Agency (JAXA), including 14 images for each of the following frame/path numbers: 7130/71, 7140/72, 7140/73, and 7140/74 (Table S1). The ALOS-1 mission operated between 2006 and 2011, carrying several sensors including PALSAR-1, a synthetic aperture radar operating at L-band with a wavelength of 23.5 cm. All images were acquired at the finest available resolution, in two polarization modes: Fine Beam Single (FBS) with HH polarization, and Fine Beam Dual (FBD) with HH and HV polarizations. The nominal spatial resolution after pre-processing was 25 m, and each scene covers a swath of approximately 70 km. All images were obtained from the Alaska Satellite Facility (vertex.daac.asf.alaska.edu/), at the Radiometric Terrain Corrected (RTC) processing level (ASF DAAC, 2015).

Since the ALOS/PALSAR mission did not perform regular observations of the Earth surface, unlike systems such as the well-known Landsat series, PALSAR images were selected from the historical record to maximize the range of flood amplitudes observed for our study site, based on the association between water stage heights observed in the Cachoeira Morena gauging station and image acquisition dates.

2.3. Image processing and classification

We adapted the approach previously used to classify Amazonian *várzea* environments, developed by Arnesen et al. (2013), Ferreira-Ferreira et al. (2014) and Silva et al. (2010). To adapt it for *igapó* forests, we had to consider the presence of a different floodplain structure, with

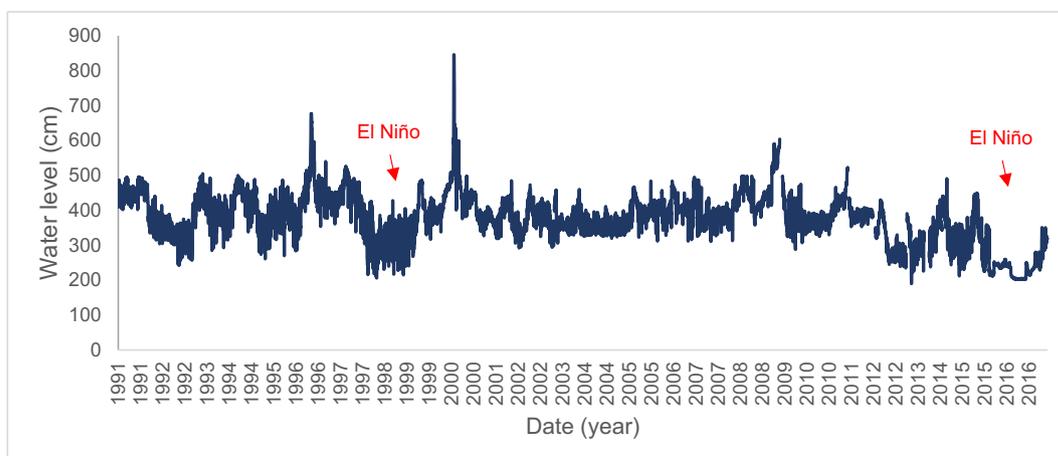


Fig. 3. The line shows the average water level during the post-dam period.

the occurrence of almost monospecific stands in lower topographies, capable of withstanding long periods of flooding (Junk et al., 2015).

Our approach is based on object-based image analysis (OBIA), where the image is first segmented into homogeneous groups of pixels (objects), which correspond to different elements of the landscape (Blaschke, 2010). Prior to segmentation, three seasonal descriptor images were calculated: (1) TBM - temporal backscattering mean, (2) TSD - temporal standard deviation of backscattering, and (3) BLW - backscattering at the lowest water level (Arnesen et al., 2013). These descriptors aid in the segregation of important landscape characteristics during segmentation, and they separate areas where there is less temporal variation from areas that change more in response to water level variation. The descriptors were converted to an 8-bit scale and filtered 3 times by a 3×3 window of a Gamma filter (Shi and Fung, 1994) to reduce speckle noise and increase the efficiency of the segmentation algorithm.

We used these three seasonal descriptors as inputs for the multi-resolution segmentation algorithm from Benz et al. (2004), iteratively optimizing the algorithm parameters as scale = 70, shape = 0.1 and compactness = 0.5, to better represent homogeneous landscape units. After segmentation, we computed the mean and standard deviation of SAR backscatter for each object in each acquired image and for the three original untransformed and unfiltered seasonal descriptors (Fig. 4).

To train the classification algorithm, we used high-resolution images from Google Earth™ and Bing Maps™, loaded into the QGIS 2.18 software using the *OpenLayers Plugin*, combined with field GPS points and field knowledge, to select about 80 training samples for each of the five pre-defined land cover classes (Table 1).

We then used the sampled training objects as inputs to the Random Forests (RF) classification algorithm, implemented in the *Random Forest* package of the open source software R (Liaw and Wiener, 2002; R Core Team, 2017). This algorithm categorizes the segments using a “forest” of

predictive trees, based on independent randomized samples (Breiman, 2001). The RF algorithm was applied following Ferreira-Ferreira et al. (2014), with the parameters *Number of Trees* and *Number of Variables Tried at Each Split* set as $n_{tree} = 15,000$ and $m_{try} = 20$.

One important limitation of our method is that it relies on the increased double-bounce scattering caused by flooding to separate floodplain forests from upland forests. We are therefore unable to detect former floodplain areas that are no longer subject to regular flooding, and some of the areas classified as upland might represent additional lost *igapó* regions, assuming the loss of flooding will lead to changes in forest structure and composition over time.

2.4. Forecasting future forest mortality

In addition to mapping dead forest stands, we used remote sensing information to identify threatened areas that are still responding to the effects of altered hydrological conditions, and where mortality is likely to be ongoing. These areas do not yet show the high degrees of mortality that make dead stands distinguishable from living flooded forest stands in the radar images.

We measured local maximum water levels at 28 known field locations, represented by the high-water marks left on tree trunks during the last flood (before June 2017). We then assigned observed maximum water levels to their respective mapped objects, which we then grouped into four floodplain classes: “high *igapó*”, “intermediate *igapó*”, “low *igapó*” (high and low referencing terrain height, according to the Junk et al. (2015) classification) and “submerged”, this last designation meaning areas that used to be low *igapó* but currently remain permanently flooded. From the images, we extracted backscattering coefficients from different dates and respective levels to represent each flood pulse phase: dry, low, medium and high water (Fig. 6).

We first analyzed the temporal radiometric behavior of each group of images corresponding to different periods of the hydrological

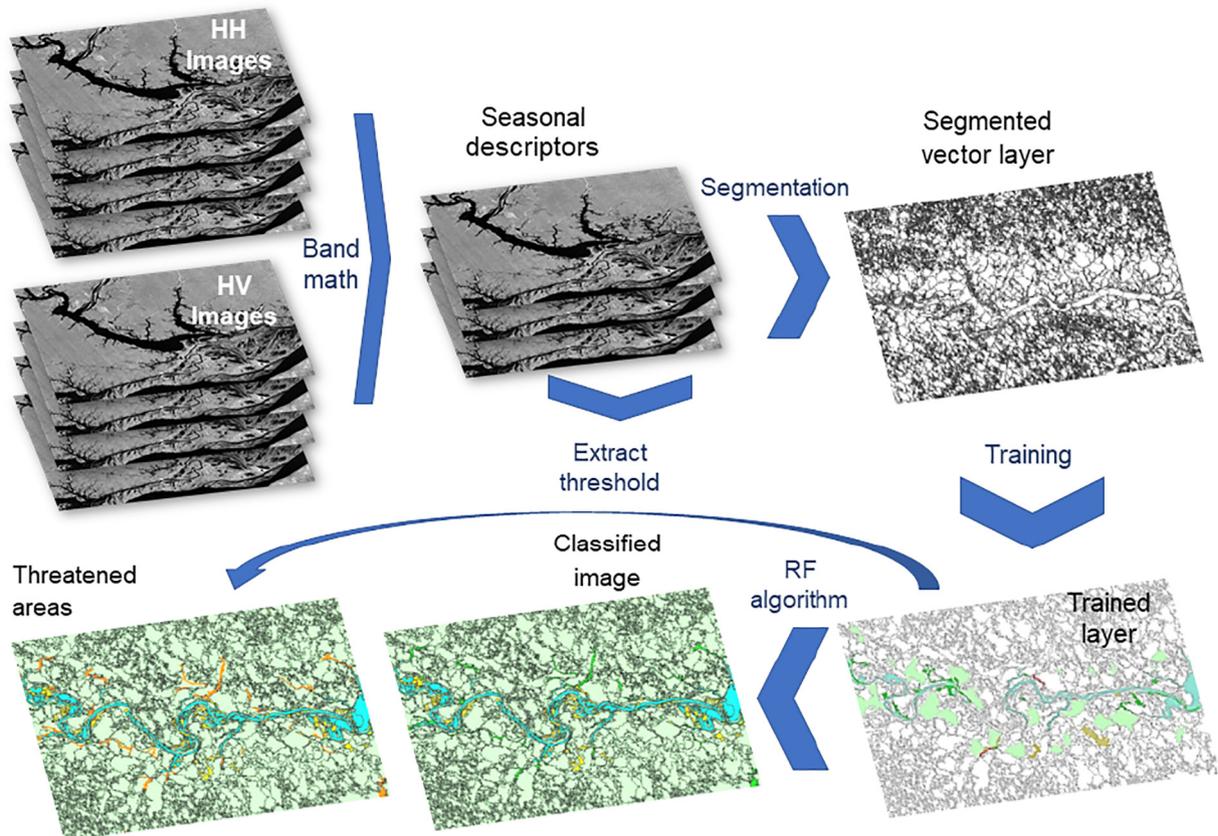


Fig. 4. Systematization of the stages of image processing, analysis and classification.



Fig. 5. Dead forest stands in the floodplains of the Uatumã River, below the blue line. Behind the dead stands, between the orange and blue lines, the igapó forest is observable, followed by upland (terra firme) forest above the orange line. (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

(flood-pulsing) regime for each floodplain class, to determine the best configuration to get good separability of classes based on the backscattered signal. Using an average image ($n = 6$) of the medium water level images (between 400 and 449 cm at Cachoeira Morena Station), we extracted the backscattering coefficients and applied a threshold to discriminate areas of “high”, “intermediate”, “low” and “submerged” *igapó*, this last class remaining flooded even during the average low water period. We defined a minimum mean flooding threshold of -4.48 dB to characterize areas that are now continuously flooded, and thereby likely undergoing mortality.

2.5. Validation

Classification accuracy was assessed using between 20 and 30 samples per vegetation class, based on high-resolution image interpretation and on observations made in the field in June of 2017, from a boat all along the focal area, using a GPS receiver. We applied the methods developed by Pontius and Millones (2011) and Pontius and Santacruz (2014) to verify the disagreement in class quantity and allocation, the latter subdivided into exchange and shift.

3. Results

Our classification using the random forests algorithm had an overall accuracy of 87.2% along the entire studied reach (Table S2). We mapped 9800 km² of flooded forest, 135 km² of non-forest areas and 13 km² of dead stands between the hydroelectric dam and the Uatumã river mouth (Fig. 7). Most flooded forest and water areas were located close to or within the confluence of the Uatumã River with the Amazon River. The water and upland classes were the most accurately classified, while the remaining three classes had some degree of confusion. Accuracy assessment showed the highest quantity disagreement errors as 12% for water and 6% for dead stands, respectively, while shift and exchange errors only occurred between flooded forest and non-forest classes, in small proportion (Fig. 8).

After mapping, we identified that, along the 300-km reach of the Uatumã River originally studied, massive mortality occurred mostly over a shorter reach, which we denominated the focal area (FA). The FA comprised an 80-km river reach beginning 43 km below the dam and included 90 km² of flooded forest and 50 km² of non-forest areas.

Massive mortality within the FA covered about 11 km², 12% of the original flooded forest cover. The furthest mapped dead stand was observed at 123 km below the dam, and about 39 km before the confluence with Abacate River.

We mapped an additional 19 km² of threatened areas, defined as floodplain forests that are now almost continuously flooded, corresponding to an additional 18% of the original flooded forest cover within the focal area. Before the Balbina Dam was built, these areas were flooded for four to eight months, whereas now they are flooded for almost the entire year and are thus very likely to be undergoing slow mortality. If we include the mapped areas of dead stands and threatened areas within the focal area, c.a. 29% of the original *igapó* forest along the 80-km focal area of the Uatumã River has been strongly impacted by the Balbina Dam (Fig. 9).

Table 1

Land cover classes defined and used to train the Random Forests algorithm to identify *igapó* forests and tree mortality in the Uatumã River (Amazon, Brazil) using ALOS/PALSAR data.

Land cover class	Standing water presence	Canopy presence	Other characteristics
Dead stands	Yes	No	Mostly dense, dead standing trees under long periods of flooding (Fig. 5).
Flooded forests	Yes	Yes	Living trees at varying densities.
Non-forest	Occasionally	No	Grasslands, shrublands, woodlands, settlements (small villages and farms), sand banks and river beaches, campinas (herbaceous vegetation/shrubs in sandy soil).
Open water	Yes	No	Permanently open water surfaces
Upland	No	Yes	Any type of forest vegetation with a closed canopy (including some campinaranas) that never flood, this class was defined as mask, and not further quantified, although in some places this category included former floodplain no longer subject to inundation.

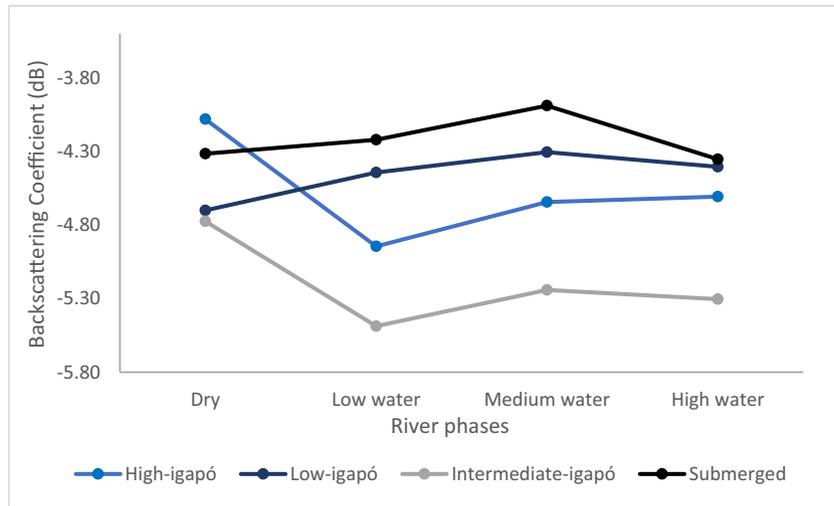


Fig. 6. L-band radar backscatter for different topographical levels of igapó forests at different periods of the flood-pulse regime.

4. Discussion

Our work shows that (1) there were significant losses of *igapó* forest downstream of the Balbina dam, caused by the changes in the flood pulse, and several areas are still undergoing mortality and may

disappear in the near future; and (2) *igapó* forests are highly sensitive to hydrological disturbances linked to loss of flood pulse seasonality; (3) the downstream impacts caused by the dams are not trivial and need to be taken into account in environmental impact assessments and reports; (4) to preserve downstream forests, the maintenance of a

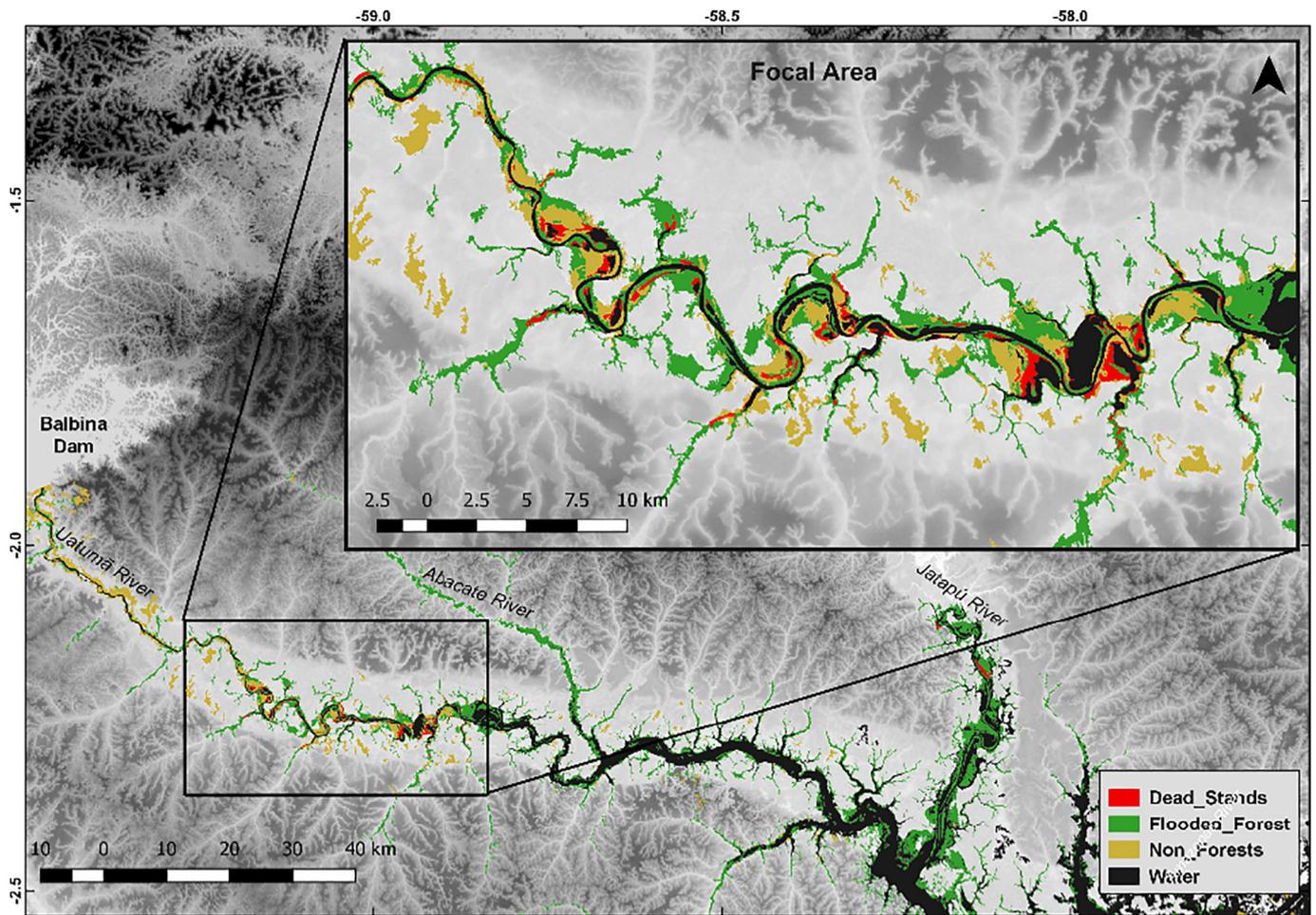


Fig. 7. Classification of the floodplains of the Uatumã River and its major tributaries (Abacate and Jatapú rivers) downstream of the Balbina dam. Dead forests occurred extensively along a stretch of about 80 km (focal area), between 43 and 123 km downstream of the Balbina hydroelectric power plant. Background: shaded relief derived from the SRTM digital elevation model.

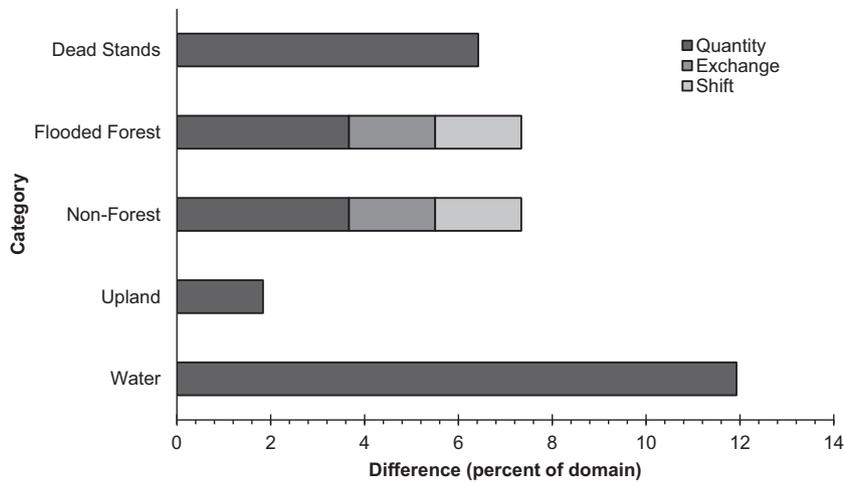


Fig. 8. Per-class disagreement between produced maps and ground-truth observations, following Pontius and Millones (2011) and Pontius and Santacruz (2014).

flood pulse that is the closest possible to the natural pulse in necessary as a mitigating action for Amazon hydroelectric projects, even if it results in reduced hydroelectric efficiency.

In the first 40 km immediately after the dam (and up to about 10 km downstream of the rapids of Cachoeira Morena), the Uatumã River follows a steep downslope along a narrow channel, where few flooded *igapós* can develop, explaining the absence of mapped floodplain forests and consequently dead stands. In addition, this reach is greatly altered by human settlements, with different land cover types and little natural

vegetation left. The effects of flood pulse changes caused by Balbina are also less evident after 120 km downstream of the dam, where the abnormal discharge caused by damming seems to be attenuated by contributions from the two largest tributary rivers (Abacate and Jatapú Rivers) and, at some point, by the Amazon River backwater effect (Meade et al., 1991).

Massive mortality occurred along regions where specific hydrogeomorphological conditions favored the formation of large floodplain forests controlled exclusively by the Uatumã River flood

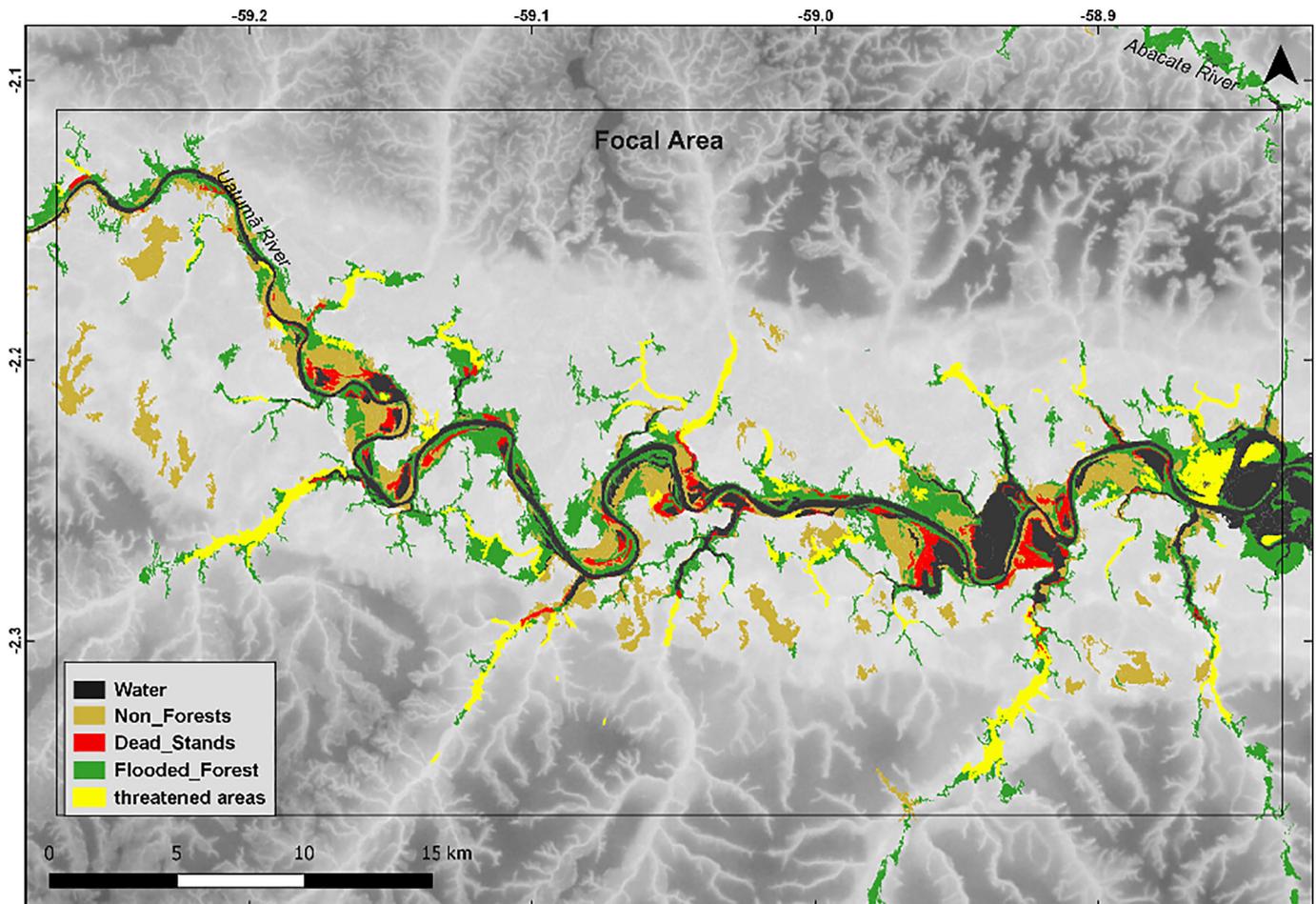


Fig. 9. Map shows potentially threatened areas adjacent to dead stands. These areas can spend entire years flooded, depending on the management of river flow by the hydropower facility.

pulse, corresponding to the reach starting 43 km after the dam and ending before the Abacate River mouth. Also, almost all dead stands were located on the insides of channel meanders, further emphasizing the connection between channel morphology and ecological impacts. Our findings corroborate Assahira et al. (2017), who made field observations of dead trees for about 100 km below the Balbina Dam.

It is important to note that our method only detects dead forest stands and potential ongoing mortality areas when they are located below current maximum flooding levels, and thus still exhibit some double-bounce response in the image series. Despite comprising a third of all flooded forest areas, these classes may not represent the entirety of lost *igapó* forests downstream; there are several past floodplain areas which are now topographically above the current maximum flooding heights and are no longer annually flooded, nor do they remain flooded for a predictable time. This change will very likely lead to profound changes in species composition, community structure and ecosystem processes over time, as the forest transitions to upland forest structure and dynamics, and thus represents a yet to be quantified additional amount of *igapó* forest loss (Assahira et al., 2017; Lobo et al., 2019).

High flood tolerance is also not a protection from the effects of flood pulse changes. In black-water *igapó* floodplains, a few tree species have the necessary adaptations to tolerate very long periods of flooding (up to 300 days per year), forming monodominant or even monospecific stands. One of the most frequent tree species in this category is *Eschweilera tenuifolia* (O. Berg) Miers, locally known as *Macacarecuia* (Junk et al., 2015; Mori and Prance, 1990). Despite being highly flood-tolerant, we observed several populations of this species suffering massive mortality within the studied area. Massive mortality was also observed for populations of the less frequent species *Macrobolium acaciifolium* within the same or similar areas, at slightly higher elevations of the Uatumã floodplains (Assahira et al., 2017).

Furthermore, a compound effect of the mortality of established adult plants is that changes in the hydrological cycle may create a barrier to seed dispersal and germination, which are generally synchronized with the flood pulse (Oliveira-Wittmann et al., 2010). Oxygen limitation in soil due to uninterrupted flooding can also lead to high mortality of potentially regenerating soil seeds and young saplings (Yang and Li, 2013), as establishment and recruitment occurs when the soil is exposed by the receding water levels (Parolin et al., 2010).

In the Central Amazonian floodplains, multiannual extreme wet and dry periods can also have dramatic effects, such as the 1971–75 period when a large-scale dieback of shrubs and trees colonizing the lowest elevations of *várzeas* and *igapós* was observed (Piedade et al., 2012). Junk (1989) related this phenomenon to a multiannual period of abnormally high minimum water levels caused by several La Niña events (Marengo, 2009; Schöngart and Junk, 2007). During strong El Niño events, the black-water *igapós* along the Negro River and other blackwater rivers are also especially vulnerable to wildfires (Flores et al., 2014; Resende et al., 2014; Sombroek, 2001), as the El Niño-induced rainfall anomalies coincide with the low-water periods of these rivers, with strong implications for plant-water availability (Schöngart et al., 2017). Large-scale fires were reported for the El Niño-induced droughts in 1925/26 (Williams et al., 2005), 1997/98 (Flores et al., 2014; Nelson, 2001) and 2015/16 (Aragão et al., 2018). These climatic vulnerabilities are expected to become more frequent in the future and may further compound the effects of river damming on floodplain forests.

Igapó ecosystems are more susceptible to wildfire than uplands (*terra firme*) or *várzeas*, due to the scarcity of nutrients and the long-term flooding that leads to an accumulation of root mats and litter on the surface, and to a drier microclimate when they are not flooded (de Almeida et al., 2016; Dos Santos and Nelson, 2013; Flores et al., 2014; Resende et al., 2014). The large amounts of dead biomass resulting from mass mortality induced by hydroelectric damming, associated with the drier microclimate in these forests, could make these ecosystems even more susceptible to wildfires (de Almeida et al., 2016;

Cochrane et al., 1999; Flores et al., 2017; Kauffman et al., 1988; Resende et al., 2014). Therefore, rising temperatures and increasing drought frequency and intensity (Gloor et al., 2013, 2015; Marengo and Espinoza, 2016) combined with land-use changes associated with deforestation and fire (Aragão et al., 2018; Nobre et al., 2016), as well as the implementation of construction of additional hydroelectric dams (Forsberg et al., 2017; Lees et al., 2016), may severely impact the *igapó* ecosystems in the future.

Massive mortality will also have several impacts on the entire ecosystem, such as niche and diversity losses, changes in community structure, facilitation of pests and diseases, and changes in biogeochemical cycles (Cherubini et al., 2002; Franklin et al., 1987). Even edge-related tree mortality on upland forests (Mesquita et al., 1999) can become a problem if the “buffer” provided by *igapó* forests is broken and adjacent upland forests are exposed. For the lowest-lying *igapós* of the Uatumã River, the carbon stock of aboveground biomass has been estimated at $118 \pm 12.8 \text{ Mg C ha}^{-1}$ (Neves, 2018). Considering the area of dead forests calculated in the present study as a conservative estimate (10.9 km^2), we can estimate a biomass loss of $129 \pm 14 \times 10^9 \text{ g C}$. However, as the total mortality process is expected to affect 30 km^2 of forests (29% of the original *igapó*), biomass loss may reach $354 \pm 38 \times 10^9 \text{ g}$. This would further worsen Balbina’s greenhouse gas emissions budget, estimated at $2.5 \times 10^{12} \text{ g C}$ (CO_2 equivalent emissions) due to increased methane (CH_4) emissions from the reservoir area (Fearnside, 1995; Kemenes et al., 2011).

In summary, our study confirms that threats to floodplains downstream of hydropower enterprises are evident, and the environmental effects of existing and proposed hydroelectric dams need to be better understood in Amazonia and other wetlands at risk (Junk et al., 2014; Kingsford, 2000; Zahar et al., 2008; Ziegler et al., 2013). Furthermore, this information must be considered when licensing and quantifying the environmental impacts of these enterprises. The legal instrument that assists the Brazilian government in making decisions about large infrastructure projects is the Environmental Impact Assessment (EIA; *Estudo de Impacto Ambiental* in Portuguese), and its accompanying Environmental Impacts Report (RIMA, in Portuguese), which must characterize the environmental conditions of an area prior to the approval of a proposed project, and must indicate all potential impacts, as well as proposing mitigating measures when appropriate (Brasil, 1986). However, the Brazilian Amazon EIA/RIMAs are often disregarded during decision making (Cochrane et al., 2017; Ritter et al., 2017). Our suggestion is to improve their diagnostic and predictive power, complementing the impact reports with surveys downstream of affected areas, as well as dynamic projective analyses of changes and predicted flood-pulse pattern after damming.

Although we could map the most fragile areas, it is still difficult to predict the exact duration and extent of the ongoing tree mortality process due to the alteration of the flood pulse. It is likely, however, that it will last for decades to come (Assahira et al., 2017) and that climate change and land-use interactions will probably intensify the deleterious effects caused by dams in the Amazon (Nobre et al., 2016). Fearnside (2016) raised the following question: “Tropical dams: to build or not to build?” Our study adds to previous evidence that the negative environmental and social effects of tropical hydropower dams are enormous, and that these effects can reach distant areas, yet they can be overlooked by scientific studies and not considered in EIA/RIMAs. This framework needs to change, and the available science must be considered to support decision-making.

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Appendix A. Supplementary data

Supplementary data to this article can be found online at <https://doi.org/10.1016/j.scitotenv.2018.12.208>.

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